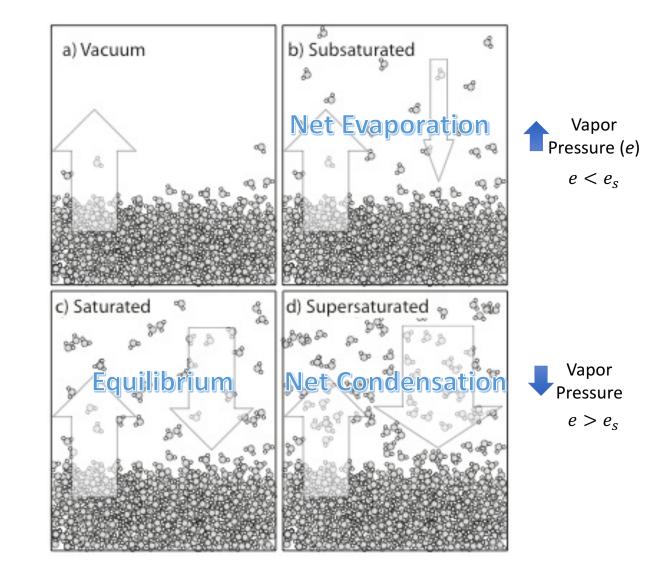


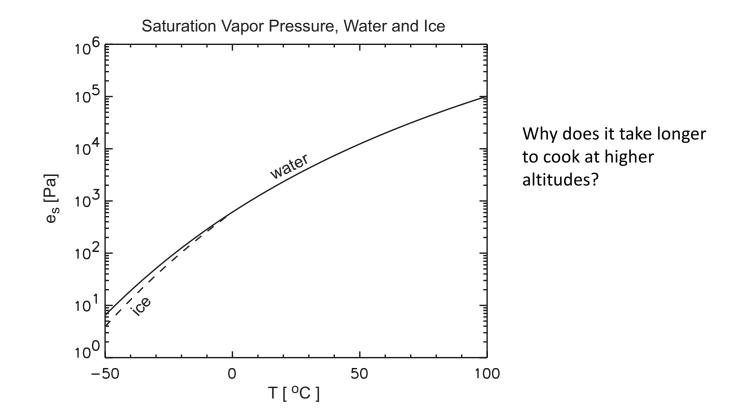
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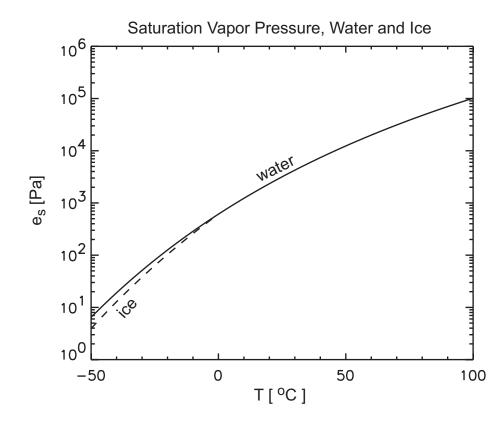
Lecture 10

- Moist Processes Part 1
 - Relative Humidity
 - Dew Point
 - Clausius Clapeyron Equation
 - Latent Heat
 - Gibbs Energy



$$e = e_s$$





Why does it take longer to cook at higher altitudes?

Lower boiling point results in less heat.

General Rule: For foods that cook in 20 minutes or less at sea level, add 1 minute of cooking time for each 1,000 feet (310 meters) of elevation.

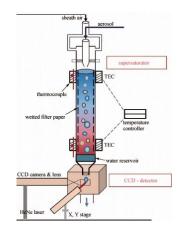
For items taking more than 20 minutes to cook, add 2 minutes for each 1,00 feet of elevation.

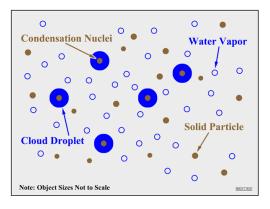
Relative humidity $RH = \frac{e}{e_s(T)}$ Supersaturation $e > e_s(T)$

 $S \equiv RH - 100\%$

 $S \approx 0.5 - 2\%$ in convective updrafts

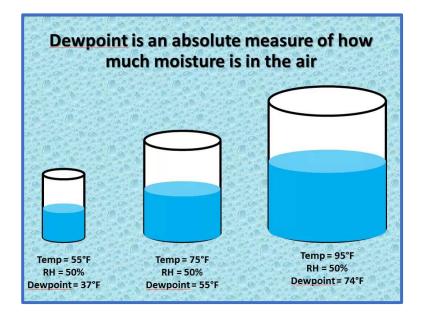






Dew Point

$$e_s(T_d) = e$$

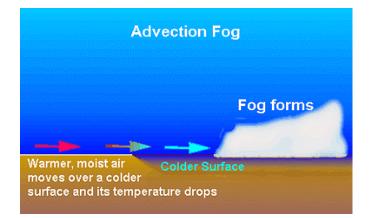


Advection Fog



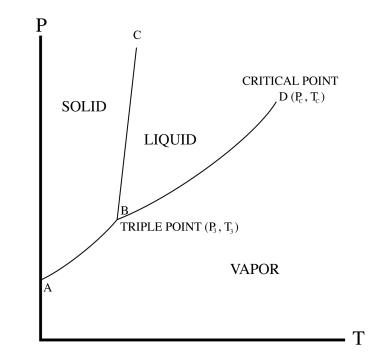
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Warm humid air moves over cold water, chilling the air to its dewpoint.



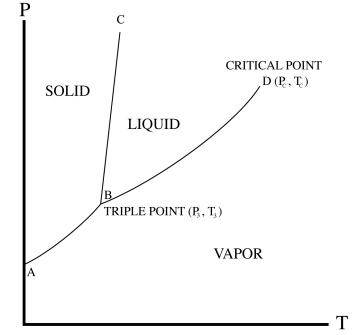
The Clausius-Clapeyron Equation

 Used to find the relationship between pressure and temperature along phase boundaries.

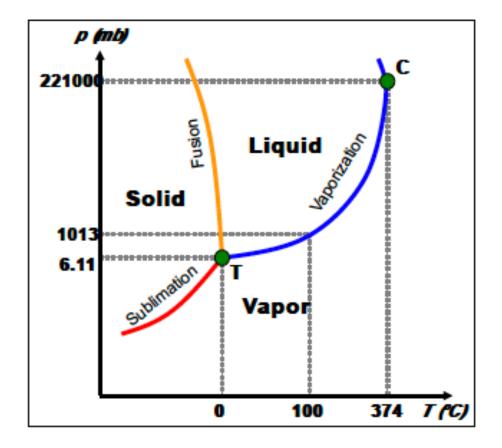


The Clausius-Clapeyron Equation

- Equilibrium vapor pressure the pressure of the vapor phase in equilibrium with a condensed (solid or liquid) phase (Line from A-D)
- Boiling temperature the temperature at which the equilibrium vapor pressure of a liquid equals the external atmospheric pressure (Line from B-D)
- Critical temperature the temperature above which there is no liquid-vapor phase transition: **Point D**, referred to as the *critical point*.
 - For every homogeneous substance, there is one *critical point*, with P_c and T_c unique to the substance.
- Melting temperature the temperature at which, for a given pressure, the solid and liquid phases coexist in equilibrium (Line from B to C)
- Triple point a fixed point in P T space corresponding to equilibrium between three phases: Point B (P₃,T₃)

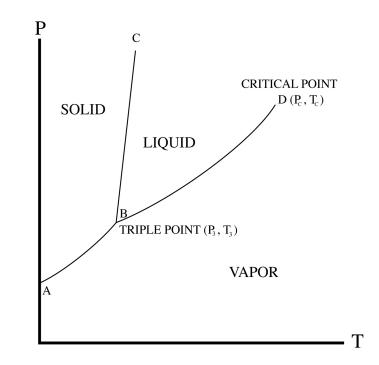


Clausius Clapeyron Diagram for H₂O



The Clausius-Clapeyron Equation

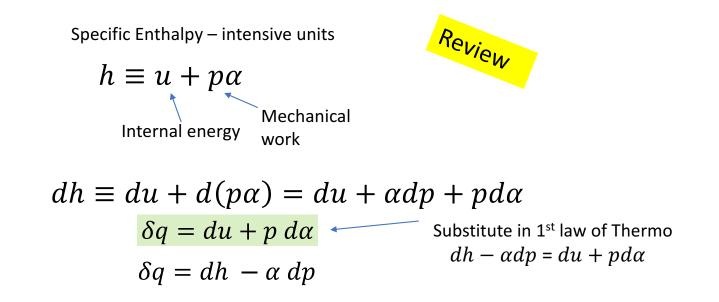
- Used to find the relationship between pressure and temperature along phase boundaries.
- Need to derive:
 - Latent Heat
 - Gibbs Energy

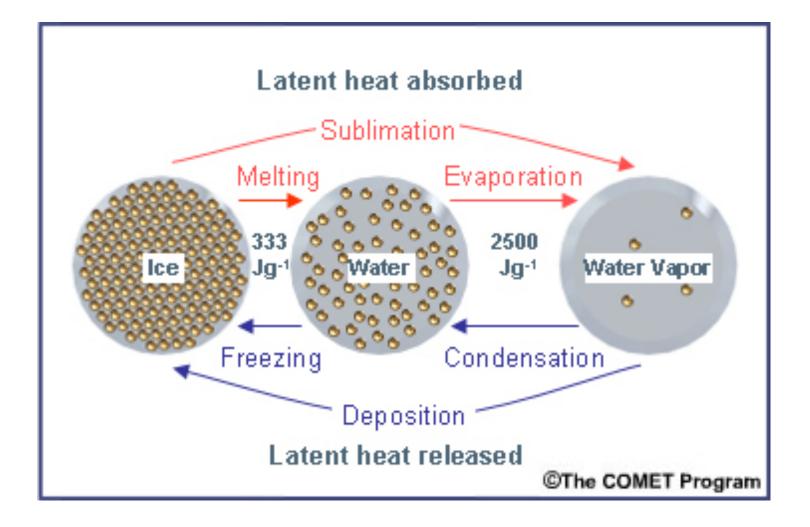


Latent Heat

- Total energy required to convert a unit mass of a substance from one phase to another while keeping the pressure and temperature constant is called the *specific enthalpy* of the phase change.
- More commonly called Latent Heat

Enthalpy – Sensible Heat





Latent Heat Specific Enthalpy at constant pressure and temperature

Specific Enthalpy - intensive units

 $h \equiv u + p\alpha$ Internal energy Mechanical work Constant Pressure $L = dh \equiv du + d(p\alpha) = \int \delta q = \int_{u_1}^{u_2} du + \int_{\alpha_1}^{\alpha_2} p d\alpha$ $p = e_s$ $L = (u_2 - u_1) + e_s(\alpha_2 - \alpha_1)$

Specific Entropy

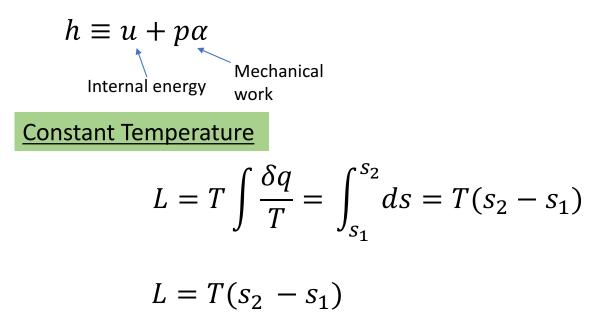
NEW STATE VARIABLE $\oint \frac{1}{T} \delta q = \oint \frac{1}{T} c_v dT + \oint \frac{1}{T} p \ d\alpha$ $\oint \frac{1}{T} \delta q = c_v \oint \frac{dT}{T} + R \oint \frac{1}{\alpha} \ d\alpha$

$$\oint \frac{1}{T} \delta q = 0 \equiv ds$$

Latent Heat

Specific Enthalpy at constant pressure and temperature

Specific Enthalpy – intensive units



Gibbs Energy $L = (u_2 - u_1) + e_s(\alpha_2 - \alpha_1)$ $L = T(s_2 - s_1)$

Combine

$$(u_2 - u_1) + e_s(\alpha_2 - \alpha_1) = T(s_2 - s_1)$$

Rearrange

$$u_1 + e_s(\alpha_1) - T(s_1) = u_2 + e_s(\alpha_2) - T(s_2)$$

Gibbs Energy = $G = u + e_s(\alpha) - T(s)$

$$G_1 = G_2$$

Remember Gibbs energy is not constant for temperature and pressure

Gibbs Energy

Gibbs Energy = G =
$$u + e_s(\alpha) - T(s)$$

dG = $du + e_s(d\alpha) + \alpha de_s - T(ds) - sdT$
Recall: $du + e_s d\alpha = \delta q = Tds$

$$dG = du + e_s(d\alpha) + \alpha de_s - (du + e_s d\alpha) - sdT$$
$$dG = \alpha de_s - sdT$$
$$G_1 = G_2$$

 $\alpha_1 de_s - s_1 dT = \alpha_2 de_s - s_2 dT$

Latent Heat Equation in 2 forms

- Remember Gibbs energy is not constant for temperature and pressure.
- Gibbs energy is the same for both the liquid and vapor phases when the two are at equilibrium (same pressure and temperature)

Clausius – Clapeyron equation

$$\alpha_1 de_s - s_1 dT = \alpha_2 de_s - s_2 dT$$

$$\alpha_1 \frac{de_s}{dT} - s_1 = \alpha_2 \frac{de_s}{dT} - s_2$$

$$\frac{de_s}{dT} = \frac{s_1 - s_2}{\alpha_2 - \alpha_1}$$
Recall: $L = T(s_2 - s_1)$

$$\frac{de_s}{dT} = \frac{s_1 - s_2}{\alpha_2 - \alpha_1} = \frac{L}{T(\alpha_2 - \alpha_1)}$$

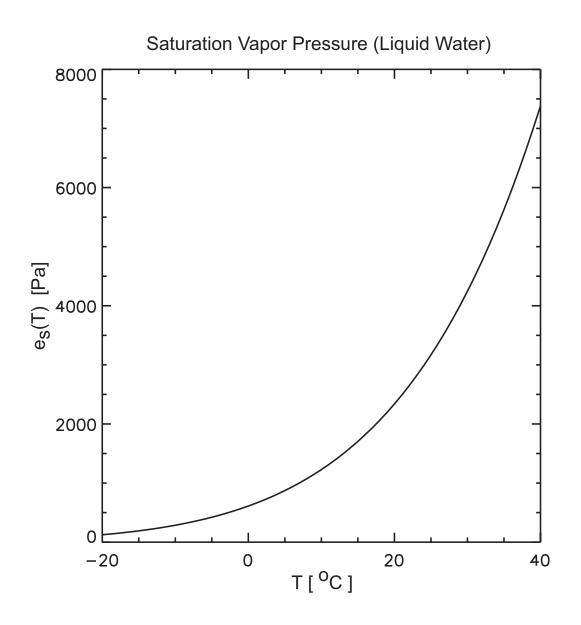
Clausius – Clapeyron equation for the atmosphere

$$\frac{de_s}{dT} = \frac{s_1 - s_2}{\alpha_2 - \alpha_1} = \frac{L}{T(\alpha_2 - \alpha_1)}$$
Assumption: $\alpha_2 \gg \alpha_1$

$$\frac{de_s}{dT} = \frac{L}{T(\alpha_2)}$$
Substitute in the ideal gas law for water vapor
$$\frac{1}{\alpha_2} = \frac{e_s}{R_v T}$$

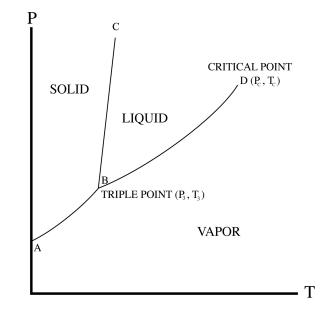
$$\frac{de_s}{dT} = \frac{Le_s}{R_v T^2} \quad \bigoplus \quad \frac{de_s}{e_s} = \frac{L}{R_v} * \frac{dT}{T^2}$$

$$\int_{e_{s0}}^{e_s} \frac{de_s}{e_s} = \frac{L}{R_v} \int_{T_0}^{T} \frac{dT}{T^2} \quad \bigoplus \quad e_s(T) = e_{s0} \exp\left[\frac{L}{R_v} \left(\frac{1}{T_0} - \frac{1}{T}\right)\right]$$



Vapor Pressure with respect to ice

$$e_i(T) = e_{i0} \exp\left[\frac{L_s}{R_v} \left(\frac{1}{T_0} - \frac{1}{T}\right)\right]$$



Supercooled water!

At T< 0 C $e_i(T) < e_s(T)$

So, at subfreezing temperature, an environment that is **saturated** with respect to water will be **supersaturated** with respect to ice.

Conversely, an environment that is saturated with respect to ice will be subsaturated with respect to supercooled liquid water

